·W. Larcher

Physiological Plant Ecology

Translated by M. A. Biederman-Thorson

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With 152 Figures

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Preface

Ecology is the science of the relationships between living organisms and their environment. It is concerned with the web of interactions involved in the circulation of matter and the flow of energy that makes possible life on earth, and with the adaptations of organisms to the conditions under which they survive. Given the multitude of diverse organisms, the plant ecologist focuses upon the plants, investigating the influence of environmental factors on the character of the vegetation and the behavior of the individual plant species.

Plant ecophysiology, a discipline within plant ecology, is concerned fundamentally with the physiology of plants as it is modified by fluctuating external influences. The aim of this book is to convey the conceptual framework upon which this discipline is based, to offer insights into the basic mechanisms and interactions within the system "plant and environment", and to present examples of current problems in this rapidly developing area. Among the topics discussed are the vital processes of plants, their metabolism and energy transformations as they are affected by environmental factors, and the ability of these organisms to adapt to such factors. It is assumed that the reader has a background in the fundamentals of plant physiology; the physiological bases of the phenomena of interest will be mentioned only to the extent necessary for an understanding of the ecological relationships. A real understanding of plant ecology requires familiarity with methodological problems and their solution; the texts by Sestak *et al.* (1971) on measurement of productivity parameters and Slavik (1974) on methods in the area of water relations are recommended.

Ecology is very much a modern field, but by no means a recent innovation. I have tried to portray this rich historical background in the choice of illustrations and tabular material; the results presented reflect the broadness of vision, the struggles and the successes of the pioneering experimental ecologists in the first half of this century, as well as the advances in knowledge made most recently.

For this English edition, the original German text has been revised, and in certain places expanded and corrected. Where possible, the German abbreviations have been replaced by those predominant in the English literature. Figures 3, 8 and 89 (formerly Fig. 90) have been extended, and three new figures have been added (Figs. 86, 145 and 152).

My first thanks are due to Dr. K. F. Springer; his publication of this English edition has made the textbook accessible to a wider circle of readers. I am grateful to the publisher of the original German edition, Roland Ulmer, for his cooperation. In particular, I thank Dr. Marguerite Biederman-Thorson for her thoughtful and sympathetic translation into English of the German text.

Above all, however, I should like to express my thanks to the pioneers of experimental ecology—Arthur Pisek, Otto Stocker, Heinrich Walter, and the late Bruno Huber. They inspired my enthusiasm for this difficult, but so attractive, field, and allowed me to benefit from their experience.

Innsbruck, September 1975

W. LARCHER

Abbreviations, Symbols and Conversion Factors

\boldsymbol{A}	Area	ϵ_p	Efficiency of radiant energy
Acc	Acceptor molecule		conversion in terms of pro-
ADP	Adenosine diphosphate		ductivity of the vegetation
ATP	Adenosine triphosphate	ϵ_F	Efficiency of radiant energy
\boldsymbol{B}	Plant biomass (also called		conversion in terms of pho-
	phytomass, the mass of a		tosynthesis
	stand of plants)	E_{p}	Evaporative power of the air;
ΔB	Change in biomass (positive	. μ	potential evaporation
	for a growing stand)	erg	Unit of energy or work
bar	Unit of pressure	U	$(1 \text{ erg} = 1 \text{ dyn} \cdot \text{cm})$
	(1 bar = 1 megadyne per	Φ	Flux, mass flow
	cm ²)	$\boldsymbol{\mathit{F}}$	Photosynthesis
C	Concentration	$F_{g'}$	Rate of gross photosynthesis
C_a	Concentration of CO ₂ and	F_n^s	Rate of net photosynthesis
-	H ₂ O in the air outside a leaf	g	Gram; unit of mass
C_i	Concentration of CO ₂ and	G	Grazing (loss of dry matter to
	H ₂ O in the intercellular sy-		consumers)
	stem of a leaf	GAP	Glyceraldehyde-3-phosphate
cal	Calorie, a unit of energy	h	Hour; unit of time
	(1 cal = 4.1868 joule =	ha	Hectare; unit of area
	$4.1868 \cdot 10^7 \text{ erg}$		$(1 \text{ ha} = 10^4 \text{ m}^2)$
D	Molecular diffusion coeffi-	hv	The energy of a light quan-
	cient		tum
d	day as a unit of time	I	Irradiance; the radiation flux
d	diameter	-	at a given level within a stand
DL_{50}	Drought lethality (degree of		of plants or body of water
	dryness causing 50% injury)	I_0	Maximum radiation flux; that
DM	Dry matter	10	incident upon a stand of
dm^2	Unit of area; for leaves, it re-		plants or body of water
	fers to one (projected) surfa-		
	ce	I_a	Long-wavelength radiation
dm_2^2	Unit of leaf area referring to	_	from the atmosphere
	the entire surface (upper and	I_{abs}	Absorbed radiation
	lower)	I_d	Direct solar radiation
E^{-}	Transpiration	I_{g}	Long-wavelength thermal ra-
E_c	Cuticular transpiration		diation from ground and
E_s	Stomatal transpiration		plants
	9		

I_i	Diffuse skylight and cloud-	LAI	Leaf-area index
	light	Lx	Lux; photometric unit of light
I_{K}	Compensation light intensity		intensity
	(at which $F = R$)	m	Meter; unit of length
\bar{I}_l	Long-wavelength radiation	M_{abs}	Quantity of minerals ab-
	balance		sorbed
I_r	Reflected short-wavelength	M_B	Mineral content of a stand of
	radiation		plants
\bar{I}_s	Short-wavelength radiation	M_G	Loss of minerals via grazing
	balance	M_{t}	Quantity of minerals incorpo-
$I_{\mathcal{S}}$	Light intensity at which pho-	3.3	rated
	tosynthesis is saturated	M_L	Loss of minerals as detritus
IR	Infrared radiation	M_r	Minerals lost in inorganic
1940	(> 750 nm)		form ("recretion")
J	Joule; unit of energy	mg	Milligram (1 mg = 10^{-3} g)
	$(1 J = 10^7 \text{ erg})$	min	Minute; unit of time
k	Coefficient, conversion fac-	ml	Milliliter (1 ml = 10^{-3} l =
	tor		1 cm ³)
k_F	Photosynthetic coefficient	mm	Millimeter; unit of length
k_{M}	Turnover factor for mineral		$(1 \text{ mm} = 10^{-3} \text{ m})$ and meas-
	nutrients in a stand of		ure of precipitation (1 mm
	plants		precipitation = 1 liter
k_{p}	Productivity coefficient		water \cdot m ⁻² of ground)
k_T	Reaction rate of biochemical	μm	Micrometer
	processes at a given tempera-		$(1 \mu m = 10^{-6} m)$
	ture	n	Number of particles
kcal	Kilocalorie	NAD+	Nicotinamide-adenine-dinuc-
	$(1 \text{ kcal} = 10^3 \text{ cal})$		leotide, reduced form:
kg	Kilogram; unit of mass		NADH + H ⁺ (simplified no-
kJ	Kilojoule (1 kJ = 10^3 joule)		tation NADH ₂); reduction
kLx	$Kilolux (1 kLx = 10^3 lux)$	NADD+	system Nicotinamide-adenine-dinuc-
kW	Kilowatt (1 kW = 10^3 watt)	NADP	leotide-phosphate, reduced
1	Liter; unit of volume		form: NADPH + H ⁺
L	Loss of organic dry matter as		(ATLEDETT) 1
	detritus		
$L_{\scriptscriptstyle E}$	Water loss via evapotranspi-	NAR	stem Net assimilation rate
	ration	nm	Nanometer Nanometer
L_I	Water loss via interception	11111	$(1 \text{ nm} = 10^{-9} \text{ m})$
L_o	Water loss via runoff and per-	OxAc	Oxalacetate
	colation	P	Turgor pressure
λ	Latent heat of vaporization	P	Production (or productivity)
	of water		of vegetation

P_{g}	Gross productivity	r_{i}	Diffusion resistances in the
P_i	Inorganic phosphate		intercellular system
P_n	Net productivity	r_p	Diffusion resistances in pro-
π	Osmotic pressure		toplasm
PEP	Phosphoenol pyruvate	r_s	Stomatal diffusion resistance
%	Percent (parts per hundred)	r_w	Diffusion resistances in the
PGA	3-phosphoglyceric acid		cell wall
pН	Negative logarithm of the hydrogen ion concentration	r_x	Carboxylation (excitation) resistance
PhAR	Photosynthetically active ra-	R	Gas constant
1 111 111	diation (400–700 nm)		$(R = 8.3 \text{ J} \cdot \text{deg}^{-1} \cdot \text{mol}^{-1})$
ppm	Parts per million	R	Respiration
Pr	Precipitation (total falling on	R_d	Dark respiration
1 /	a stand of plants)	R_{l}	Respiration in the light
Pr_n	Precipitation reaching the	RH	Relative humidity
1 ' n	ground beneath a plant can-	RuDP	Ribulose-1,5-diphosphate
	opy	RuP	Ribulose-5-phosphate
Ψ	Water potential	S	Second; unit of time
PWP	Permanent wilting percen-	t	Time (point in time or dura-
IVVI	tage	L 1	tion)
Py	Pyruvate	t	Ton (Metric; $1 t = 10^3 \text{ kg}$)
Q	Energy flow	T	Temperature (all temperature
Q_E	Energy conversion associated	141	data in °C)
£.E	with evaporation and conden-	T_a	Air temperature
	sation	T_a T_l	Leaf temperature
Q_H	Energy conversion associated	au	Matric pressure or potential
\mathcal{L}_H	with convection	TL ₅₀	Temperature-stress lethality
Q_I	Energy conversion associated	1 L ₅₀	(the temperature at which
ΣI	with radiation from the sun		50% of plants are killed by
	and reradiation		heat or cold)
Q_{M}	Energy conversion associated	9 Se 2007	(40)
\mathcal{L}_{M}	with metabolism	torr	Unit of pressure (1 torr =
Q_P	Energy conversion in plant		$1.33 \cdot 10^{-3} \text{ bar} \doteq \text{a 1-mm co-}$
\mathcal{L}_{P}	communities	* * * *	lumn of Hg)
0	Energy conversion in the	UV	Ultraviolet radiation
Q_{Soil}	soil		(< 400 nm)
0		W	Watt; unit of power
Q_{10}	Temperature coefficient of biochemical and physiolo-	•••	$(1 W = 1 J \cdot s^{-1})$
	gical process	W	Weight
r	Transport or diffusion resi-	W_{abs}	Quantity of water absorbed
r	stance	W_{act}	Actual water content (when
	Boundary (air) layer resis-		sample is taken)
r_a	tance	W_{av}	Available water
		av	

W_d	Dry weight	WSD	Water saturation deficit
W_{FC}	Water content of soil at field	yr	Year
8	capacity	Z	Relative height or depth
W_{PWP}	Water content of soil at per-	Ø	Diameter
	manent wilting percentage	÷	Approximately equal to
W_f	Fresh weight	>	Larger than
W_s	Water content in saturated	<	Smaller than
	state		

Equivalents

Radiation

$$\begin{array}{l} 1 \ cal \cdot cm^{-2} \cdot min^{-1} = 6.98 \cdot 10^{5} \ erg \cdot cm^{-2} \cdot s^{-1} = 6.98 \cdot 10^{-2} \ W \cdot cm^{-2} \\ 1 \ erg \cdot cm^{-2} \cdot s^{-1} = 1.43 \cdot 10^{-6} \ cal \cdot cm^{-2} \cdot min^{-1} = 10^{-7} \ W \cdot cm^{-2} \\ 1 \ W \cdot cm^{-2} = 10^{7} \ erg \cdot cm^{-2} \cdot s^{-1} = 14.3 \ cal \cdot cm^{-2} \cdot min^{-1} \\ 1 \ cal \cdot cm^{-2} \cdot min^{-1} \doteq 60 \ kLx \\ 100 \ kLx \doteq 1.5 \ cal \cdot cm^{-2} \cdot min^{-1} \end{array}$$

Pressure

1 bar =
$$10^6$$
 dyne \cdot cm⁻² = 1.019 atm (technical)
1 atm (technical) = 1 kg-weight \cdot cm⁻² \doteq 0.981 bar

Energy Consumption in the Evaporation of Water

Heat of vaporization at
$$0^{\circ}$$
 C = 597 cal \cdot g⁻¹ H₂O at 10° C = 592 cal \cdot g⁻¹ at 20° C = 586 cal \cdot g⁻¹ at 30° C = 580 cal \cdot g⁻¹

Phytomass

1 g DM · m⁻² =
$$10^{-2}$$
 t · ha⁻¹
1 t · ha⁻¹ = 100 g · m⁻²
1 g org. DM $\doteq 0.45$ g C $\doteq 1.5$ g CO₂
1 g C $\doteq 2.2$ g org. DM $\doteq 2.7$ g CO₂
1 g CO₂ $\doteq 0.67$ g org. DM $\doteq 0.37$ g C

Further aids to conversion can be found in the *Manuals of Methods* by Šestak *et al.* (1971) and Slavik (1974).

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The Environment of Plants

Plants have colonized nearly all regions of the earth, including the oceans and inland waters; on land they can be found even in such inhospitable places as deserts and fields of ice. Far back in geological time, when the first land plants were evolving, they encountered a world of water, air and stone. That is, their environment consisted of the hydrosphere, atmosphere and lithosphere. Later, as the cover of vegetation gradually closed, and with the assistance of microorganisms and animals, there developed the most important substrate of plants: the soil—the pedosphere.

The Hydrosphere

The hydrosphere comprises the oceans of the world, which cover an impressive 71% of the earth's surface, as well as the inland waters and the groundwater. Great differences exist in the chemical compositions of these bodies of water (Fig. 1). Sea water, rich in Na⁺, Mg²⁺, Cl⁻ and SO₄² and with an average salt content of 35 g·l⁻¹, differs fundamentally from fresh water, which usually contains more Ca²⁺ and HCO₃⁻; but there are local differences as well, depending on the nature of the inflowing waters and the degree of mixing. Moreover, currents have an effect upon temperature gradients. Where there are no currents, the strong absorption of radiation in the upper levels of the water leads to a characteristic layering with respect to temperature and density; this has a marked influence upon nutrition, productivity and distribution of aquatic organisms (see p. 19).

The Atmosphere

The air enveloping the earth provides plants with carbon dioxide and oxygen. It also mediates the balance of water through the processes of rain, condensation and "evapotranspiration". Continual movement of the air ensures that its composition remains fairly constant—79% nitrogen (by volume), 21% oxygen and 0.03% carbon dioxide, water vapor and noble gases (Fig. 1). In addition the air contains gaseous, liquid and solid impurities; these are primarily sulfur dioxide, unstable nitrogen compounds, halogen compounds, dust and soot.

The part of the atmosphere with which plants come into contact is the troposphere, the weather zone of the earth's envelope of air. The nature of this zone varies over short distances and is characterized in several ways: (1) by the weather (short-term events such as showers, thunderstorms and gusts of wind), (2) by meteorological events of intermediate duration such as periods of rain or frost and (3) by the climate (the average state and ordinary long-term fluctuations in meteorological factors at a given place). Depending on the terrain and on the density, height and type of vegeta-

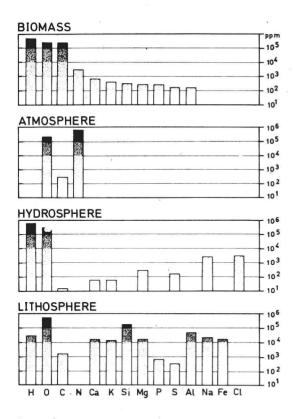


Fig. 1. Composition of the biomass, atmosphere, hydrosphere and lithosphere, in terms of the relative numbers of atoms (atoms per million atoms, not the proportion by weight) of the various chemical elements. The composition of living organisms is clearly distinct from that of the three components of their environment: they select from the available elements, according to their needs. The scale of the ordinate is logarithmic. For example, in the biomass H, O, C and N are present in the greatest proportions: 4.98 · 105 atoms per million (i. e., about 50% of all atoms) are hydrogen atoms: oxygen and carbon atoms each comprise 24.9 · 105 atoms per million (about 25%), and 2.7. 103 (about 0.3%) are nitrogen atoms. (After Deevey, 1970)

tion, individual climatic regions of different sizes are formed. Within the large-scale "macroclimate" measured by the network of meteorological stations, one may distinguish "microclimates" that prevail in specific places such as certain slopes, narrow valleys and stands of vegetation; an "interface" climate—in the layer of air near the ground and the surface of leaves—may also be distinguished. Thus the parts of plants above ground are exposed to variability, in space and time, with respect to radiation, temperature, humidity, precipitation and air motion; any of these can from time to time represent a threat to the organism.

The Lithosphere and the Soil

The earth's crust is the inexhaustible storehouse for the variety of chemical elements of which organisms are composed (Fig. 1). The lithosphere exchanges matter with the hydrosphere, and also affects the composition of the atmosphere through volcanic action and the products of radioactive decay. Primarily, however, it is the basic material for the formation of the soil.

Soil is more than just superficially loosened lithosphere. It is the product of the transformation and mingling of mineral and organic substances. Soils are produced with the assistance of organisms and under the influence of environmental condi-

tions. They are subject to continual change; soils grow, mature, and can age and perish. Physical and chemical weathering continuously frees mineral substances from the rocky substrate, and there is an unceasing decay of plant remains and dead organisms. These decay products, together with the excrement of soil animals, gradually turn into humus, which forms complexes with the mineral products of weathering. In natural soils a profile is established of more or less horizontal layers ("horizons"); between the strata of litter and humus, and the stratum where weathering of the parent rock occurs, there are transitional zones with varying proportions of humus. The types and thicknesses of the horizons in such a profile are characteristic of a given type of soil and reflect the influence of climate, plant cover, soil organisms, underlying rock and the activity of man. Pedology is the science of soil formation and composition, and of the classification of soil types. Knowledge of the fundamentals of this field is an absolute prerequisite to understanding plant ecology. The solid particles in the ground stick together to form aggregates, leaving small open spaces. Together these form a system of pores penetrating the entire soil, filled partly with air and partly with water. Thus the soil is a three-phase system in which lithosphere (the solid phase), hydrosphere (fluids) and atmosphere (gases) are intermingled. It has an enormous capacity for uptake and storage, and is particularly suited for the buffering of physical and chemical influences. Below the top centimeters of soil the prevailing climate is more stable than that of the atmosphere; radiation is essentially unable to penetrate, there are no sharp gradients of temperature, and the processes of exchange are slow, occurring by diffusion. Therefore the soil is the most suitable habitat for many organisms. The roots—in many respects the most vulnerable organ system of the higher plants—are entirely adapted to life in the soil. A landscape without soil is a life-repelling "lunar" landscape. Only a few remarkable

The Biosphere

sand.

Atmosphere, hydrosphere and lithosphere existed before there was life on earth, but it is of course only through the appearance of living organisms that they became significant as an "environment". "Environment", as defined by A. F. Thienemann, is "the totality of external conditions affecting a living organism or a community (biocenosis) of organisms in its habitat (its biotope)". In this strict sense, only living beings have an environment. It comprises not only the influences exerted by the abiotic surroundings, but also those due to the other organisms present.

plants such as aerial algae, lichens and mosses can actually thrive on bare stone or

The part of the earth which supports life, called the biosphere, is the narrow band, about 100 m thick, above the earth and below the surface of the ocean, which is ordinarily inhabited by organisms. It is true that birds can fly as high as 2,000 m, and that there are bacteria on the floor of 10,000-m-deep marine trenches, but abundant life is limited to a much more restricted region near the surface of the earth. Trees stand no more than 70–100 m tall and sink their roots but a few tens of meters into the earth; in water the layer penetrated by light (and hence densely populated) ordinarily extends to a depth of 30 m, and at most to 100 m.

Among living organisms, the plants are of prime significance. They are capable of capturing and storing by photosynthesis the energy from outer space—that of sunlight—and in terms of mass they far exceed all other organisms; about 99% of the total mass of living beings (the biomass) on earth is accounted for by the plants (the phytomass). Because of this enormous mass, the plant cover is a stabilizing factor in the cycling of matter and has a crucial effect upon the climate.

The Ecosystem: Interplay of Biological and Environmental Factors

Communities of organisms and their abiotic surroundings interact in nature by many kinds of reciprocal relationships, both structural and functional. Circumscribed, more or less uniform sections of the biosphere delimit biogeocenoses, or ecosystems. The expression "biogeocenosis" was coined by V.N. Sukachev, and the term "ecosystem" is attributable to R. Woltereck and A.G. Tansley. H. Ellenberg defines the ecosystem as "a unitary system of interactions involving living organisms and their inorganic environment which is, to a certain degree, capable of self-regulation". In short,

Each ecosystem has a certain spatial extent; together ecosystems form a diverse mosaic in the biosphere. A forest is an ecosystem, as is a meadow, a lake or an ocean. The principles by which all these ecosystems operate hold equally well for the biosphere as a whole, and as well for natural ecosystems as for artificial systems such as aquariums and self-contained manned spacecraft. Fig. 2 represents the typical structure of an ecosystem and the most important interactions both within it and between it and the external world.

The Components of an Ecosystem

Every independently functional ecosystem is composed of at least two biological components: the producers and the decomposers. Between these there may exist a whole chain of consumers.

Primary Producers are the autotrophic organisms, which effect the incorporation of inorganic elements in organic compounds and thus raise them to a higher energy level. The green plants and some bacteria utilize sunlight to form carbohydrates from carbon dioxide and water (photosynthesis), and these become the basis for further syntheses. Various microorganisms accomplish the same thing by using the energy freed by exergonic inorganic reactions (chemosynthesis).

Consumers (or phagotrophs) are the heterotrophic organisms that feed directly or indirectly upon the organic substances synthesized by the primary producers. The principal consumers are the herbivores and the plant parasites. The herbivores serve in turn as food for the carnivores, and both are attacked by animal parasites.

Decomposers (or saprotrophs) are those organisms that finally reduce plant and animal refuse to the level of its basic inorganic components. This group includes

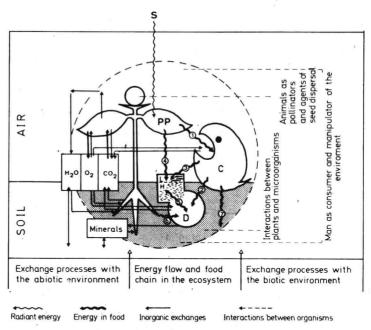


Fig. 2. The flow of energy and matter, and the cycling and exchange of substances, in a terrestrial ecosystem. Components of the ecosystem: PP primary producers; C consumers; D decomposers; L deposits of detritus (plant litter, bodies of animals); H humus. Energy flow and the food chain; S radiation from the sun; 1 consumption of plants as fodder and by parasites; 2 organic excretions from animals and microorganisms; 3 detritus consisting of animal cadavers and dead microorganisms; 4 detritus derived from the primary producers; 5 decomposition (humification and mineralization) of detritus; 6 organic excretions from the plants; 7 loss of organic waste from the ecosystem. Inorganic transport of substances: CO_2 to primary producers (photosynthesis) and from processes of breakdown in primary producers, consumers and decomposers (soil respiration), O_2 from primary producers (photosynthesis) to oxygen-consuming catabolic processes; H_2O from evaporative surfaces (from the ground and through organisms) into the atmosphere, as precipitation from the atmosphere into the soil, from the soil through consumption by organisms and loss by drainage; mineral substances from the soil into primary producers and back again via the food chain and the activity of decomposers (mineralization)

bacteria and some soil animals. The decomposers, like the herbivores and other consumers, can serve as food for other organisms. In such cases, they take on the role of secondary producers. Thus a single individual, depending on its position in the food chain, can play the role of secondary producer, consumer or decomposer.

Food Chains and Energy Flow in an Ecosystem

A "food chain" is defined as the sequence of stages through which stored energy in the form of food is passed from the primary producers to a number of organisms. As