



Hydrology and Water Resources Engineering

K.C. Patra

江苏工业学院图书馆 藏 书 章



Alpha Science International Ltd.

0018528

Kanhu Charan Patra

Associate Professor Department of Civil Engineering College of Engineering and Technology O.U.A.T., Bhubaneswar-751 003, India

Copyright © 2001

Alpha Science International Ltd. P.O. Box 4067, Pangbourne RG8 8UT, UK

Exclusive distribution in North America only by CRC Press LLC

All rights reserved. No part of this publication may be reproduced, stored in a retrieval system or transmitted in any form or by any means, electronic, mechanical, photocopying, recording or otherwise, without the prior written permission of the publishers.

ISBN 1-84265-023-8

Printed in India.

To the memories of those engineers and scientists, who have enhanced our knowledge on water resources and hydraulic engineering

Water is precious for life on earth. There is an ever increasing demand for the supply of fresh water to the various sectors of the human needs. This has given rise to the problem of optimal management of water resources potential on all parts of the world, more so, in a developing country like India, where the distribution of the resource is highly uneven both in space and time. Hydrology plays the central role in the development and management of water resources and therefore, the protection of environment. As a result, hydrology forms a part of curricula at the undergraduate and postgraduate levels in agriculture, civil engineering, environmental engineering, geology and earth sciences, forestry and meteorology.

Based on my experience as an engineer in the water resources design, planning, development and management for more than a decade and a long teaching experience thereafter, I felt the need for a textbook written in simple and lucid style that can be easily followed by students who have not had an opportunity of the exposure to the subject before. A student can understand most parts of this book himself. At the same time, the book can serve as a source of information to the engineers and other professionals dealing with water resources planning, development and management. Therefore, the chapters have been designed and developed in a systematic way. The materials for the book are mostly taken from the class notes prepared for the graduate students and at the same time utilising my decade long experience from the field problems. Therefore, emphasis have been given on the applicability of the text materials to the field situations. Using data from the field, a large number of hydrologic design problems have been worked out for each concept of the chapters to illustrate the analysis and design procedure.

The text is arranged in ten chapters. Chapter 1, besides a brief resume to the history of hydrology, introduction to meteorology, discusses hydrologic cycle, cloud and availability of water on earth. Since statistics and probabilities play an important role in the estimation of hydrologic parameters, the same have been discussed in Chapter 2. Chapter 3 is essentially devoted to the precipitation, its measurement and analysis. Measurement and estimation of hydrologic losses like evaporation, evapotranspiration, interception, depression storage and infiltration are covered in detail in Chapter 4. Groundwater forms a major component in hydrologic cycle and therefore, the hydrologic aspects of ground water are discussed in Chapter 5. In Chapter 6, the aspects of stream flow measurement and the rainfall runoff process are discussed at length. Hydrographs are discussed in Chapter 7. Estimation of design flood for various types of catchments is given in Chapter 8, while Chapter 9 covers various methods of channel and reservoir routing. In Chapter 10, reservoir sedimentation and the aspects of cost benefit have been discussed. Sufficient problems are given at the end of each chapter for practice. Important references have been cited at the end of the book.

The author is grateful to Er. B.B. Singhsamant (Chief Engineer) and Er. N. C. Mohanty (Executive Engineer), Water Resources Department, Government of Orissa, for their active help and continuous encouragement while preparing the manuscript of this book. I am also to record my sincere thanks to Dr. S.K. Kar, Professor, Department of Civil Engineering, IIT Kharagpur and Dr. S.C. Mishra, Dean, College of Engineering, for their active help and support. I am also grateful to all the authors, whose works I have consulted.

Finally I thank my wife Gita and our children Sonam and Suman for their tolerance and understanding during the preparation of the manuscript.

I am thankful to N.K. Das and R. Mohan for their help in typing the manuscript and drawing figures in Autocad.

Any suggestion in improving the book will be thankfully received and will be incorporated in the next edition.

K.C. PATRA

CONTENTS

D			
Prefe	ice		vii
	T .	Dependeng Models 78 ST Name 1	
1.		roduction of a laborate purposite parameter	1
	1.1	General 1 See and to washing the real purposed at	
	1.2	History of Hydrology 1	
	1.3	Meteorology 3	
		Lapse Rate 4 Pressure 5	
		Water Vapour 5	
		Precipitable Water 7	
		Latent Heat 9	
		Humidity 10	
	1.4	Cloud and Raindrop Formation 10	
	1.5	Hydrologic Cycle 11	
	1.6	Availability of Water on Earth 14	
	1.7	Importance of Hydrology and its Applications in	
		Engineering 16	
		Problems 17	
2	Stat	istics and Probabilities in Hydrology	10
4.	2.1	istics and Probabilities in Hydrology Introduction 19	19
	2.1		
	2.2	Statistical Parameters 20 Central Tendency Parameters 20	
		Dispersion Characteristics 25	
		Skewness 26	
		Kurtosis 27	
	2.3	Theoretical Probability Distribution 30	
		Discrete Distribution 31	
	2.4	Continuous Distribution 33	
	2.4	Frequency Analysis 38 Gumbel's Distribution 40	
		Pearson Type III Distribution 41	
		Log Pearson Type III Distribution 41	
		Normal Distribution 42	
	0.5	Log Normal Distribution 43	
	2.5	Graphical Method Using Probability Paper 50	
		Construction of Probability Paper 51 Selection of Type of Distribution 51	
	2.6	Confidence Limits 52	
	2.7	Regression and Correlation 56	
	2.7	Graphical Method 56	
		Analytical Method 56	
		Correlation 58	
		Significance of Parameters 60	
		Standard Forms of Bivariate Equations 61	

	2.8	Multivariate Linear Regression and Correlation 64 Multiple Correlation Coefficient 65	
	2.9		
		Trend Analysis 70	
		Oscillation 73	
		Jump '74	
		Periodicity or Seasonality 74	
		Serial Correlation 76 Stochastic Component 78	
	2.10	Dependenc Models 78	
	2.10	TOTAL CONTRACTOR OF THE CONTRA	
	2.11	Chi-square Test of Goodness of Fit 82	
		Problems 90	
3	. Pre	cipitation	94
	3.1	Introduction 94	77
	3.2	Forms of Precipitation 95	
	3.3	Types of Precipitation 96	
	3.4	Rainfall in India 99	
	3.5	Measurement of Rainfall 101	
	3.5	Non-Recording Rain Gauge 102	
		Recording Type Rain Gauges 103	
		Weather Radar 106	
		Totalisers 107	
	3.6	Network Design 107	
		Optimum Number of Rain Gauge Stations 108	
	27	Ideal Location for a Rain Gauge Station 109	
	3.7	Consistency of Rainfall Data 111	
	3.8	Estimating Missing Data 114	
		Arithmetic Mean Method 114	
		Normal Ratio Method 114 Regression Method 115	
		Inverse Distance Method 115	
	3.9	Presentation of Precipitation Data 118	
		Moving Average Curve 119	
		Mass Curve 119	
		Rainfall Hyetograph 119	
	2 10	Intensity-Duration-Frequency Curves 121	
	3.10	Mean Aerial Rainfall 122	
		Arithmetic Mean Method 123	
		Thiessen Polygon Method 123 Isohyetal Method 125	
		Grid Point Method 126	
		Orographic Method 127	
		Isopercental Method 127	
	3.11	Depth-Area-Duration Curve 129	
		Design Storm 131	
		Statistical Storm 132	
		Probable Maximum Precipitation 140	
		Standard Project Storm 145	

Los	ses from Precipitation	149
4.1	Introduction 149	
4.2	Evaporation and its Estimation 149	
	Measurement Using Evaporation Pans 151	
	Empirical Equations 154	
	Water Balance Method 156	
	Energy Budget Method 157 Mass Transfer Method 161	
4.3	Methods to Reduce Reservoir Evaporation 161	
4.4	Evapotranspiration (Consumptive Use) and its	
4.4	Estimation 163	
	Experimental Measurement 164	
	Climatic Approaches 165	
4.5	Interception 176	
4.6	Depression Storage 178	
4.7	Infiltration and its Estimation 178	
4./	Factors Affecting Infiltration 179	
	Field Measurement Using Infiltrometers 182	
	Rainfall-Runoff Analysis 184	
	Infiltration Indices 185	
	Mass Curve Method 190	
4.0	Analytical Models 191	
4.8	Watershed Leakage 191	
	Problems 192	
Cro	ound Water	195
5.1	Introduction 195	175
	7	
5.2	Zoning of Subsurface 195	
	Aeration Zone 195 Saturated Zone 197	
5.3	Water Pearing Materials 107	
5.4	Aquifer Properties Affecting Ground Water 199	
3.7	4 'C M ' 1 D ' 200	
	Ground Water Flow Parameters 202	
	Stratification of Layers 207	
	Calculation of Flow from Flownet 209	
5.5	Steady Flow Equation 211	
5.6	Unsteady Flow Equation 214	
5.7	Ground Water Flow Problems 215	
	Steady Flow in Unconfined Aquifer 215	
	Steady Flow in Confined Aquifer of Constant Thickness 217	
	Steady Flow in a Confined Aquifer of Variable Thickness 219	
	Unconfined Flow with Recharge from Rainfall 220 Drainage Using Tiles 222	
	Flow Through Leaky Aquifer 224	
	Flow into Infiltration Galleries 226	
5.8	Well Hydraulics 227	
	Steady Radial Flow into a Well 228	
	Partially Penetrated Well 233	
5.9	Unsteady Flow into a Well 235	
	Confined Aquifer 235	
	Unsteady Radial Flow into a Well in Unconfined Aquifer 242	

5.10	Spacing of Wells 243	
5.11	Well Loss 245	
5.12	Well Adjacent to a Stream (Method of Images) 246	
5.13	Sea Water Intrusion 248	
5.14	Methods of Ground Water Investigation 250	
	Hydro-Geologic Investigation 250	
	At Surface 252 Sub-Surface Geophysical Methods 253	
	Model Study 255	
	Hydrologic Investigation 255	
	Problems 256	
Stre	am Flow	25
6.1	Introduction 250	40
6.2	T II1 250	
6.3	Factors Affective Devet of 201	
6.4	Stage Measurement 266	
0.1	Non-Recording Stage Recorders 266	
	Automatic Gauge Recorders 268	
6.5	Discharge Measurement 270	
	Discharge Measuring Structures 270	
	Approximate Areá-Slope Method 270 Slope Method 273	
	Area Velocity Method 276	
	Radio Tracer Method 286	
	Dilution Technique 286	
6.6	Electromagnetic Induction Method 287 Stage Discharge Relationship 287	
0.0	Stage Discharge Relationship 287 Extension of Stage Discharge Relation 290	
6.7	Requirement of a Good Gauge-Discharge Site 295	
	Network Design 296	
6.8	Runoff Computation 297	
	Extension of Runoff Records 297	
6.9	Runoff From Rainfall Records 305	
	Use of Rainfall-Runoff Data at a Site 305	
	Use of Rainfall-Runoff Relation of Neighbouring Sites 308 Empirical Relations 310	
6.10	Runoff Simulation Models 315	
	Steps in Modelling 316	
	HEC Model 317	
	Problems 321	
Hvd	rograph	32
7.1	Introduction 324	24
7.2	Hydrograph Concept 325	
7.3	Components of Hydrograph 326	
7.4	Unit Hydrograph 334	
	Assumptions and Conditions in Unit Hydrograph 335	
	Limitations of Unit Hydrograph 336	
7.5	Uses of Unit Hydrograph 336	
7.5	Derivation of Unit Hydrograph 337	
	From Complex Storms 341	

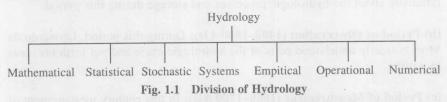
	7.6	S-Hydrograph 346	
	7.7	Change of Unit Duration of Unit Hydrograph 349	
		Required Duration is an Integer Multiple of D-hour 349	
		Required Duration is a Real Multiple of D-hour 350	
	7.8	Instantaneous Unit Hydrograph (IUH) 350	
	7.9	Derivation of UH from IUH 352	
	1.9	Derivation of Instantaneous Unit Hydrograph 353 From S-hydrograph 353	
		From Conceptual Models 353	
	7.10		
		Snyder's Approach 379	
		Synthetic Unit Hydrograph for Indian Catchment 381	
	7 11	Goel et al. Method 382	
		Dimensionless Unit Hydrograph 384	
	7.12	Distribution Graph 386 Problems 388	
		Problems 300	
8.	Desi	gn Flood	392
	8.1	Introduction 392	
	8.2	Design Flood 392	
	8.3	Flood Peak Estimations for Ungauged Catchments 395	
		Rational Method 395	
		Empirical Equations 399	
	8.4	Envelope Curves 408 Flood Estimation for Gauged Catchment 410	
	0.4	Flood Estimation for Gauged Catemnent 410	
		Unit Hydrograph Approach 420	
	8.5	Regional Flood Frequency Analysis 427	
	8.6	Analysis of Partial Duration Series 430	
		Problems 432	
0	121	1 D - C	
9.		d Routing	434
	9.1	Introduction 434	
	9.2	Routing Methods 435	
	9.3	Hydrologic Channel Routing 437	
		Muskingum Equation 439 Working Value Method 447	
	9.4	Hydraulic Channel Routing 449	
		Finite Difference Method 450	
		Numerical Methods in Routing 452	
	9.5	Hydrologic Reservoir Routing 453	
		Analytical (Trial and Error) Method 454	
		Modified Puls Method 455 Inflow-Storage-Discharge Method 461	
		Goodrich Method 466	
	9.6	Flood Routing Machines 469	
		Mechanical Flood Routers 469	
		Electric Analog Routing Machine 469	
	9.7	Digital Computers 470 Flood Forgasting 470	
		Flood Forecasting 470 Flood Control Measures 482	
	2.0	Flood Control Measures 482 Problems 485	
		110011111111111111111111111111111111111	

10.	Res	ervoir and Sedimentation	4	188
		Introduction 488		
		Fixation of Reservoir Capacity 488		
		Ripples Mass Curve 489		
	10.2	Sequent Peak Algorithm 490		
		Determination of Spillway Size 494		
	10.4	Allocation of Storage Space for Various Uses 496		
		Reservoir Elevation-Area-Capacity Curve 497 Reservoir Operation 497		
		Reservoir Working Table 498		
	10.5	Reservoir Sedimentation 501		
	10.6	Determination of Sediment Yield at a River Site 502		
		Sheet Erosion 502		
		Sediment Measurement by Sample Recorder 504 Bed Load Estimation 507		
		Empirical Relations for Total Sediment Load 509		
	10.7	Reservoir Sedimentation 513		
		Reservoir Classification 514		
	10.0	Distribution of Sediment in Reservoirs 515		
		Reservoir Sediment Control 521		
	10.9	Reservoir Economics 524		
		Cost-Benefit Ratio 524 Optimisation of Benefits 528		
		Linear Programming in Multipurpose Water Resource Projects 530		
		Problems 532		
	Овл	ECTIVE QUESTIONS	53	34
	REF	PERENCES	54	17
	Con	NVERSION FACTORS	54	
	INDI	Control of the contro		
			55	03

Introduction

1.1 GENERAL

Hydrology deals with the origin, distribution and circulation of water in different forms in land phases and atmosphere: It is an interdisciplinary subject starting with meteorology through agronomy, forestry, geology, hydraulics and finally oceanography. Statistics, physics and chemistry help to formulate and understand the subject more conceptually. Solutions to hydrologic problems are usually carried out by borrowing techniques from several disciplines. Depending on the treatment of data, hydrology can be divided into various sub-branches as shown in Fig. 1.1.



Broadly, hydrology is classified into two groups. The 'scientific hydrology' is concerned mainly with academic aspect, whereas, the 'engineering' or 'applied' hydrology includes: (i) estimation of water resources; (ii) study of transmission process like precipitation, evaporation, runoff and their interdependence; (iii) understanding the properties of water in nature and (iv) dealing with natural problems like droughts and floods. The distribution of water potential in India is highly uneven both in space and time. With the growth of population, the demand of water has also increased to many fold. The science of hydrology can be effectively employed to overcome these difficulties. Therefore, a lot of research and developmental opportunities do exist in this comparatively new branch of science. We would describe some of the aspects of hydrology, which should help a student to understand the subject. The book should also help field engineers or hydrologists to solve their problems connected with water resources engineering.

1.2 HISTORY OF HYDROLOGY

There is no evidence of the date of construction of the first water resource project on the earth. Water being the chief ingredient in life, all ancient civilisations flourished only near the sources of water and then probably collapsed either when the water supply failed or when they were devastated by floods. Man's quest for knowledge, his capacity for innovation and adventure have contributed

greatly to his efforts at understanding and exploiting this resource right from Indus-valley civilisation. Archaeological evidence of the existence of wells in Mohenjodaro (3000 BC), the diversion weirs on the river Nile by king Means (3000 BC) and the kanats for irrigation in Persia were the few structures men believed to have constructed in the ancient times. Reference can also be found in Vedic literature (3000 BC), Arthashastra by Chanakya (300 BC), Brihatsamhita by Varahamihira (505–587 AD) and in lots of other ancient writings about the availability, measurement and use of the surface and ground water. In India, during Chanakya's period, kings probably maintained a network of gauges in their kingdoms for tax collection. Writings of ancient thinkers like Aristotle (350 BC), Plato (350 BC), Vitruvius (100 BC) and others suggest their scientific thinking of various hydrologic processes. Following Chow, the history of hydrology can be classified into the following eight periods.

- (a) Period of Speculation (Prior to 1400 AD): During this period, many hydraulic structures were known to have been constructed. The important being the Arabian wells, Persian Kanats, Egyptian and Mesopotamian irrigation projects, Roman aqueducts, water supply and drainage projects in Indus valley and the Chinese irrigation systems. Vitruvius was probably the first great philosopher to think rationally about the hydrologic processes and storage during this period.
- **(b) Period of Observation (1400–1600 AD):** During this period, Leonardo da Vinci correctly understood parts of the hydrologic cycle and put forth his ideas in writing.
- (c) Period of Measurement (1600–1700 AD): In this century measurement of rainfall, runoff, evaporation and study of artesian wells were taken up by different thinkers. Correct predictions of some hydrologic phenomena were also made.
- (d) Period of Experimentation (1700–1800 AD): New discoveries like Bernoulli's piezometer, Pittot tube, Waltman's current meter, Borda's tube, Bernoulli's theorem, Chezy's formula and D'Alembert's principle were made during this period. All these helped to quantify the total water resource potential of the world.
- (e) Period of Modernisation (1800–1900 AD): Substantial contributions were made in the fields of both ground and surface water hydrology during this century. Important contributions being Darcy's law, Poiseuillie's equation of capillary flow, Dupuit-Thiem's well formula, Weir-discharge formula, Kutter's determination of Chezy's equation, Manning's equation, Price current meter and Dalton's law of vapour pressure.
- (f) Period of Empiricism (1900–1930 AD): Thirty years following the early part of 20th century saw the science of hydrology largely empirical. A number of formulae were proposed and various governments undertook massive efforts in setting up organisations for measurement of hydrological parameters quantitatively in most of the important river basins of the world.

- (g) Period of Rationalisation (1930-1950 AD): Most of the present concepts in hydrology date back to 1932, the year in which Le Ray K. Sherman rationalised the concept of hydrology by demonstrating the use of Unit Hydrograph for translating the rainfall excess into runoff hydrograph. The other great contributions during this period were Horton's (1933) determination of rainfall excess from infiltration characteristics, Theis (1935) well hydraulics, Gumbel's (1941) extreme value frequency distribution, Bernard's introduction of meteorology to hydrology and Einstein's theory of sediment load. A lot of laboratories of international repute were also established.
- (h) Period of Theorisation (Since 1950): The advent of modern high speed computing machines, sophisticated instrumentation and emergence of modern fluid mechanics helped greatly in the formulation of new theories on hydrology. Many hydrologic activities of international level were also established. Some notables are Water Resources Development Centres (WRDC), World Meteorological Organisation (WMO), International Association of Scientific Hydrology (IASH) etc. In 1961 an International Hydrologic Decade (IHD) was proposed by IASH to co-ordinate international research and training programs in hydrology.

In India, National Institute of Hydrology (NIH) and Department of Hydrology at the University of Roorkee (besides other institutions) are promoting research and training programs and co-ordinating international activities.

1.3 METEOROLOGY (METEOROS-LOFTY; LOGOS-DISCOURSE)

It is that branch of science which deals with the entire gaseous envelope around the earth. Hydrology starts with precipitation of water vapour on the earth's surface. Therefore, it is desirable to know something about the factors, which bring about these changes. Atmosphere (atoms-vapour; spheria-sphere) is the gaseous envelope of the earth. It is a mixture of gases and the water vapour is of special interest to the hydrologists. The atmosphere exerts a weight of about 1.033 kg/cm², which is equivalent to 10.3 m depth of a water column. The atmospheric pressure is one atom (~ 1013.25 mb). Though water vapour in the atmosphere is limited to a few kilometres from earth's surface, it may be required to know the division of atmosphere into concentric layers. The layers with the variation of temperature and pressure across it are shown in Fig. 1.2.

Atmosphere enveloping the earth is divided into five concentric layers. Troposphere (a layer just above the earth's surface) contains all moisture, dust and almost three-fourth of total air mass. Tropopause is the top of troposphere, which separates it from the stratosphere layer above it. The height of troposphere varies from 8 km at poles to 16 km at the equator. In the stratosphere, temperature is almost constant. Mesosphere is just above stratosphere. Temperature in this zone is always higher than ground surface due to absorption of heat from sun's radiation. The fourth layer is known as ionosphere, which is 70-80 km above the earth's surface and is responsible for reflecting ordinary radio waves. This layer has a high concentration of free electrons. The outermost shell is called exosphere, where atmosphere loses all its properties. A sound knowledge of troposphere is important for a hydrologist as the weather is confined to this layer only.

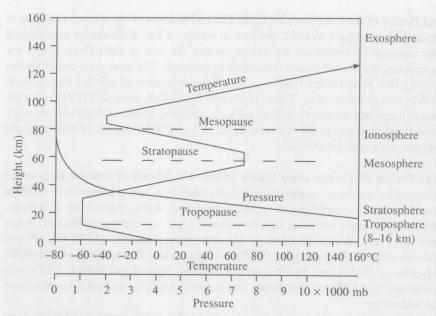


Fig. 1.2 Variation of Temperature and Pressure with Height from Earth's Surface

1.3.1 Lapse Rate

The rate at which temperature decreases with the increase of altitude at any particular time is called lapse rate. When a mass of air at the surface of earth is heated up, it expands and becomes lighter than the surroundings and then rises. Similarly, heavier air at higher altitudes sinks and contracts due to increase in pressure. Though no heat is added or subtracted from the system, the change in temperature takes place due to expansion and contraction only and this change in temperature for the dry air, known as dry adiabatic lapse rate is 10°C/km. The dry adiabatic lapse rate is higher than the normal air lapse rate which varies at the rate of 6.5°C/km up to the first 10 km from earth's surface and then almost at zero rate for the next 10 km. Inversion is the reverse of lapse rate. As air mass rises, it loses heat with altitude and at certain stage there is no further decrease in temperature even if it rises further unless there is condensation of some water vapour. This temperature at which the air mass just becomes saturated, if cooled at constant pressure without any moisture added or removed from it, is called dew point. After condensation level, the air mass continues to ascend due to the latent heat of condensation added to it. This change in temperature is called adiabatic saturation lapse rate, which is between 3 and 10°C/km of the altitude change.

Sometimes the term *normal temperature* meaning the mean of 30 records of the corresponding period picked up from each year is used. Temperature data published by India Meteorological Department (IMD) contains daily, weekly, monthly or annual normal records. For example, the normal temperature of 1st January is obtained by taking the average of 30 years of 1st January data from say 1970 to 1999. This may be either the minimum or maximum temperature for

the day. On the other hand, the term 'mean' refers to the daily averages of maximum and minimum temperatures. For example, for 1st January 1999, the average of maximum and minimum temperatures for the day is the mean temperature.

1.3.2 Pressure

Atmospheric pressure is the pressure exerted by the gases of the atmosphere on the earth's surface due to their own weight. It decreases uniformly from the surface of earth and the rate of decrease depends on temperature and lapse rate. Again due to excess heat, air at the equator belt rises and sinks at around 30° latitude belts. Therefore the equator belt is of low pressure zone and the 30° north and south belts are the higher pressure zones (sub-tropical belt). At 60° north and south, the higher pressure air coming from poles meet the air coming from sub-tropical high pressure zones and causes it to rise to form a low pressure belt at this zone. Due to earth's rotation around its own axis the movement of a parcel of air is affected (coriolis force). This force, along with the fictional resistance on the earth's surface, forces the air at an angle across the isobars causing a counter clockwise movement around low pressure centres and clockwise movement around high pressure centres in the north hemisphere. A reversal of the situation can be seen in the southern hemisphere. At equator, the speed of the earth is 1670 km/h from west to east and at 60° latitude it is about 830 km/h. From the knowledge of physics, we can calculate a theoretical velocity of 2505 km/h of a parcel of air at rest, relative to earth's surface at the equator and moving east-northward direction to 60° latitudes. However, this does not happen because of frictional resistance the air has to overcome at the surface due to features of the earth and air densities. An oversimplified pattern of air circulation on earth's surface is shown in Fig. 1.3.

Isobar (line joining points of equal pressure) is drawn on a weather map in such a way that the lines pass through points of equal sea level pressure. A front is the border region of adjacent air masses, normally, the colder air forms a wedge above which warmer air ascends. The ascending warmer air cools with the lapse rate leading to the formation of cloud and precipitation. Front is of four types. A cold front is formed at the boundary between the advancing cold air and retreating warm air (Fig. 1.4a), whereas in a warm front, the advancing warm air meets the retreating cold air (Fig. 1.4b). In a stationary front, both the warm and cold masses cease to move. When the two cold fronts meet, the region between two fronts become warmer. This situation is conducive for the formation of cyclones. Both fronts get elevated and the condition is called occluded front.

1.3.3 Water Vapour

It is that gaseous stage of water which can be condensed. The amount of water vapour in the atmosphere is a meagre 1 part in 100000 of all the water on earth but it plays a vital role for the sustenance of life on earth. A large amount of water passes through the atmosphere due to the continuous process of vaporisation and condensation taking place globally. The process by which liquid is converted to vapour is called evaporation or vaporisation. Condensation is the reverse